

# Height datum unification within a global vertical reference system

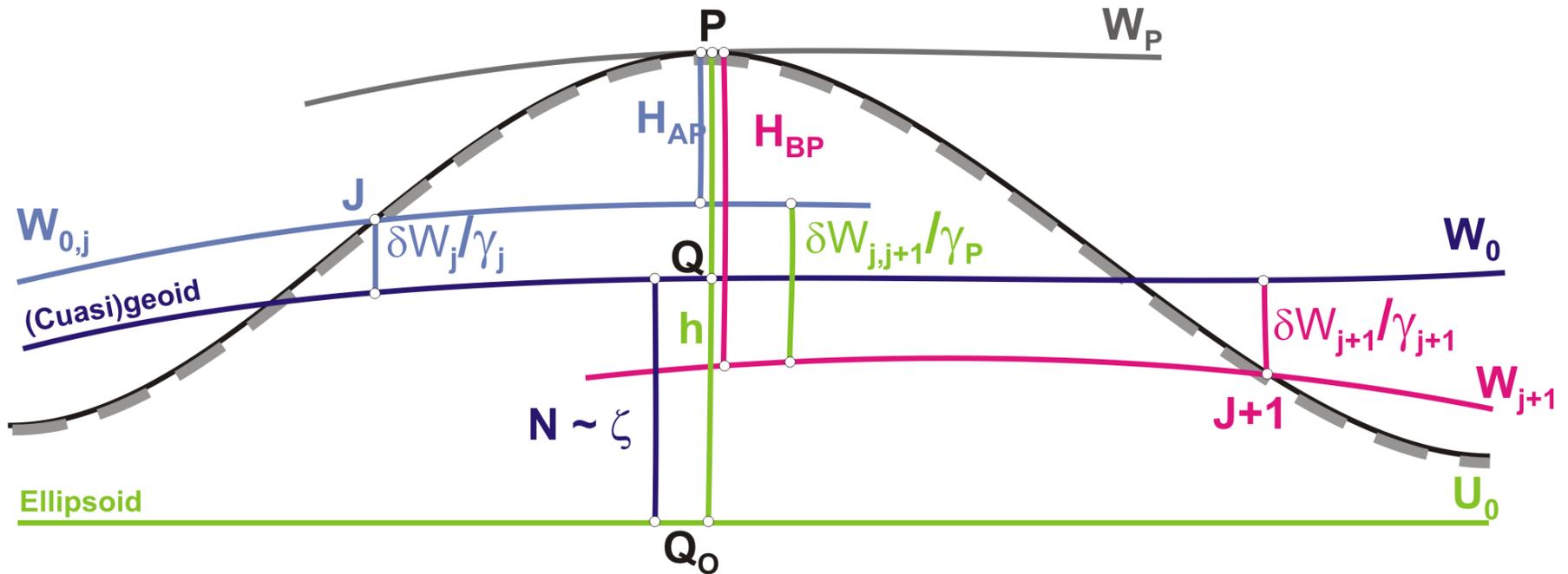


Laura Sánchez  
Deutsches Geodätisches Forschungsinstitut (DGFI)



Geodesy for Planet Earth, IAG Scientific Assembly 2009  
September 2, 2009  
Buenos Aires, Argentina

# Objectives



1. To satisfy  $H = h - \zeta$  with high precision world-wide;
2. To define and realize a global vertical reference system supporting geometric and physical heights;
3. To integrate the existing local height systems into the global one.

# Basic approach

## Definition

type of coordinates,  
reference surfaces,  
consistency between  
geometric and  
physical heights

## Realization

- Conventions to realize the definition ( $W_0$ , tide system, reference epoch, etc.)
- Establishment of a global reference frame (similar to ITRF)
- Determination of (vertical) coordinates for the reference frame according to the definition and conventions
- Unification of the existing local height systems into the global one
  - SSTop at and around reference tide gauges
  - Connection of the local levels to the ITRS/ITRF
  - Connection of neighbouring local height systems
  - Connection parameters at epoch of local level definitions
    - Time variations of sea level at the reference tide gauges
    - Separation of crustal movements from sea level changes
    - Vertical movements of height benchmarks
- Re-calculation of the height related observables and iteration of the realization procedure until getting a mm-level accuracy

# Modern definition of a vertical reference system

Consistent modelling of geometric and physical parameters, i.e.  
 $h = H^N + \zeta (\approx H + N)$  in a global frame with high accuracy ( $> 10^{-9}$ )

## Geometrical Component

Coordinates:

$h(t), dh/dt$

Definition:

**ITRS + Level ellipsoid ( $h_0 = 0$ )**

- ( $a, J_2, \omega, GM$ ) or
- ( $W_0, J_2, \omega, GM$ )

Realization:

- Related to the **ITRS** (ITRF)
- Conventional ellipsoid

Conventions:

IERS Conventions

**Ellipsoid constants,  $W_0$ ,  $U_0$  values, reference tide system have to be aligned to the physical conventions!**

## Physical Component

Coord.: Potential differences

$-\Delta W_P(t) = W_0(t) - W_P(t); d\Delta W_0/dt$

Definition:

$W_0 = \text{const.}$  (as a convention)

Realization:

- Selection of a global  $W_0$  value
- Determination of the local  $W_{0,j}$  values
- Connection of  $W_{0,j}$  with  $W_0$
- Geometrical representation of  $W_0$  and  $W_{0,j}$  (i.e. geoid comp.)
- Potential differences into physical heights ( $H$  or  $H^N$ )

**Zero tide system**

# Realization of $W_0$ ( $W_{0,j}$ ) by solving GBVP

## Ocean areas

Fixed gravimetric GBVP  
provides the global reference level ( $W_0$ )

## Land areas

Scalar-free GBVP  
provides the local reference levels ( $W_j$ )

**Formulation**

$$\nabla^2 T = 0 \quad \text{outside boundary surface} \quad ; \quad T = W - U$$

$$-\frac{\partial T}{\partial r} = \delta g$$

$$-\frac{\partial T}{\partial r} - \frac{2}{r}T = g_j - \frac{2}{r}\delta W_j$$

$$\delta W_j = W_{0,j} - U_0 = W_0 - W_{0,j}$$

**Constraints**

$$T = 0 \text{ at } \infty ; \int_{\text{sea}} T d\sigma = k_i ; \int_{\text{land}} T d\sigma = k_j ; \int k_i + \int k_j = 0 \Rightarrow T_{00} \equiv \int T d\sigma = 0$$

**Solution**

$$T = \frac{\Delta GM}{R} + \frac{R}{4\pi} \iint_{\sigma} B_j S(\psi) d\sigma + \sum_{n=1}^{\infty} \frac{R}{4\pi} \iint_{\sigma} G_n S(\psi) d\sigma$$

$$j = 1 ; B_1 = \delta g \text{ (gravity disturbances)}$$

$$j = 1 \dots J ; B_j = g_j - \frac{2}{r}\delta W_j$$

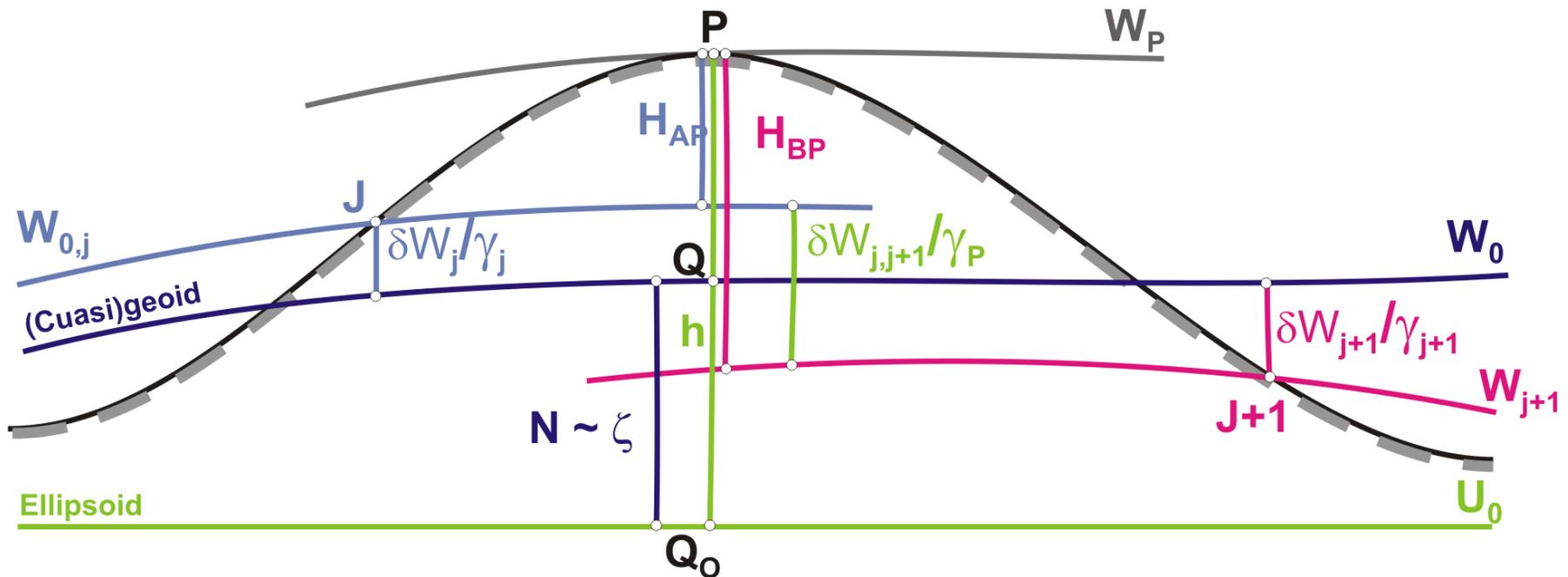
$$g_1 = \Delta g ; g_2 = \Delta C ; \text{ etc.}$$

vertical datum unification strategies

# Strategies for the vertical datum unification

A global vertical reference system has to be connected to the geometric terrestrial reference system (TRS) to satisfy  $h = H + \zeta$  world-wide. This is possible by constraining the determination of the  $\delta W_j$  terms to:

$$\gamma_P h_P - (W_0^j - W_P^j) - T_P^j - 2\delta W_j = 0$$



# Strategies for the vertical datum unification

## Oceanic approach

(SStop around gauges)  
Satellite altimetry and satellite-only GGM, SStop at coast lines including also tide gauge records.

$$T_P^j - T_0 = \delta W^j$$

## Coastal approach

(reference tide gauges)  
GPS positioning at tide gauges, (geopotential numbers), terrestrial gravity and satellite-only GGM.

$$\frac{1}{2}T_P^j - \frac{1}{2}h_P\gamma_P = \delta W^j$$

## Terrestrial approach

(geometrical reference stations)  
GPS positioning at reference stations (including border points), geopotential numbers, terrestrial gravity and satellite-only GGM.

$$\frac{1}{2}(T_P^j + \Delta C_P^j) - \frac{1}{2}h_P\gamma_P = \delta W^j$$

$$\frac{1}{2}(T_P^j + \Delta C_P^j) - \frac{1}{2}(T_P^{j+1} + \Delta C_P^{j+1}) = \delta W^{j+1} - \delta W^j$$

Solution of observation equations by a combined adjustment !

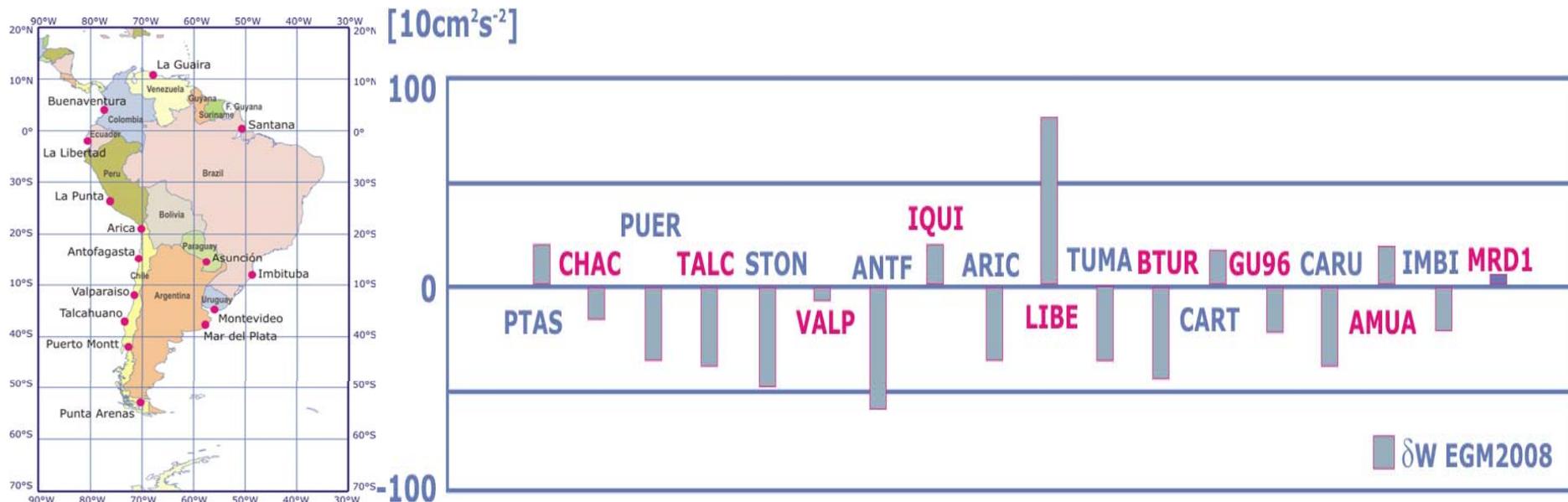
# Numerical tests: coastal approach

Solution of the fixed GBVP on ocean areas:  $W_0 = 62\,636\,853,1 \text{ m}^2 \text{ s}^{-2}$

Geometry of the mean sea surface: CLS01 model (Hernandez, Schaeffer 2001),  
Gravity disturbances from CHAMP-GRACE GGM, n=150, coefficients @ 2000.0

Solution of the scalar-free GBVP at reference tide gauges

SIRGAS coordinates, EGM2008 model (coefficients @ 2000.0), zero tide system.



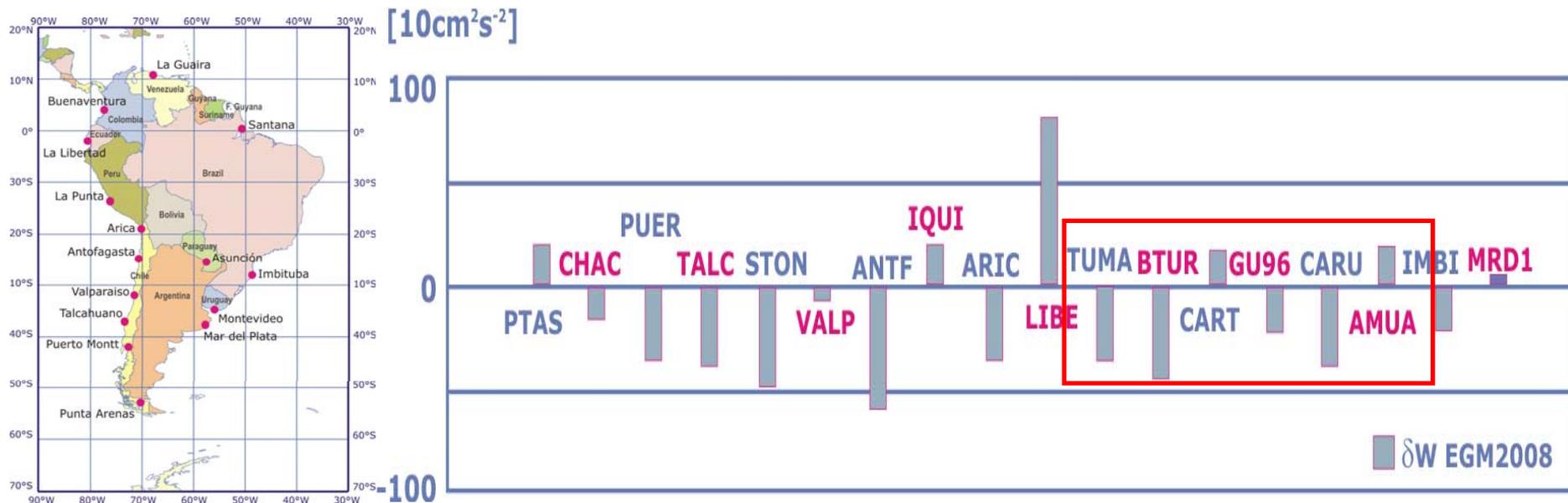
# Numerical tests: coastal approach

Solution of the fixed GBVP on ocean areas:  $W_0 = 62\,636\,853,1 \text{ m}^2 \text{ s}^{-2}$

Geometry of the mean sea surface: CLS01 model (Hernandez, Schaeffer 2001),  
Gravity disturbances from CHAMP-GRACE GGM, n=150, coefficients @ 2000.0

Solution of the scalar-free GBVP at reference tide gauges

SIRGAS coordinates, EGM2008 model (coefficients @ 2000.0), zero tide system.



# Numerical tests: coastal and oceanic approach

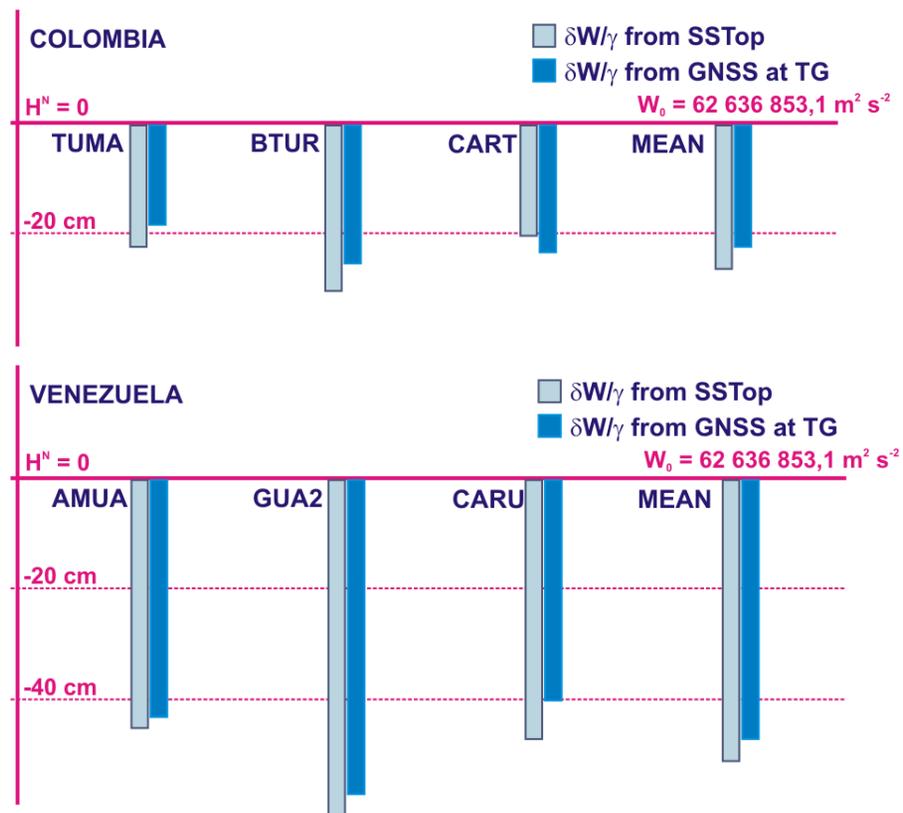
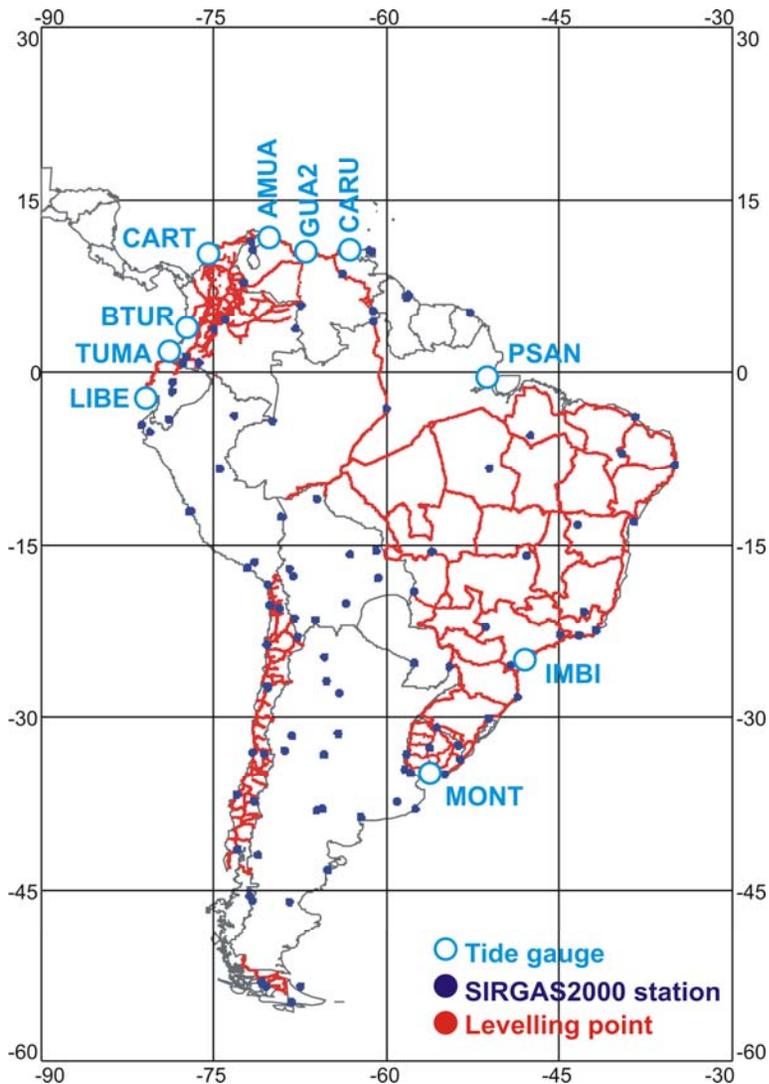
Input data:

Local quasigeoid models

GNSS positioning

Geopotential numbers from levelling

$H^N$ ,  $h$ ,  $\zeta$ ,  $SSTop$  at epoch 2000.0, zero tide system



# Outlook

- To extend the numerical tests to the entire SIRGAS region. This implies:
  - Determination of  $SSTop$  at all the reference tide gauges;
  - GNSS positioning at the same tide gauges to distinguish sea level changes from vertical movements of the Earth's crust;
  - Continental adjustment of the first order vertical networks;
- To estimate the connection terms  $\delta W_i$  at the definition period of the local reference levels, i.e. all heights ( $h$ ,  $H^N$ ,  $\zeta$ ,  $SSTop$ ) must be reduced to a common reference epoch;
- To determine the  $\delta W_i$  using regional (or local) geoids of high resolution. The GGMs do not provide the required accuracy and resolution;
- To follow the IAG recommendations about a **reliable** and **precise** global reference level (a  $W_0$  conventional value).