

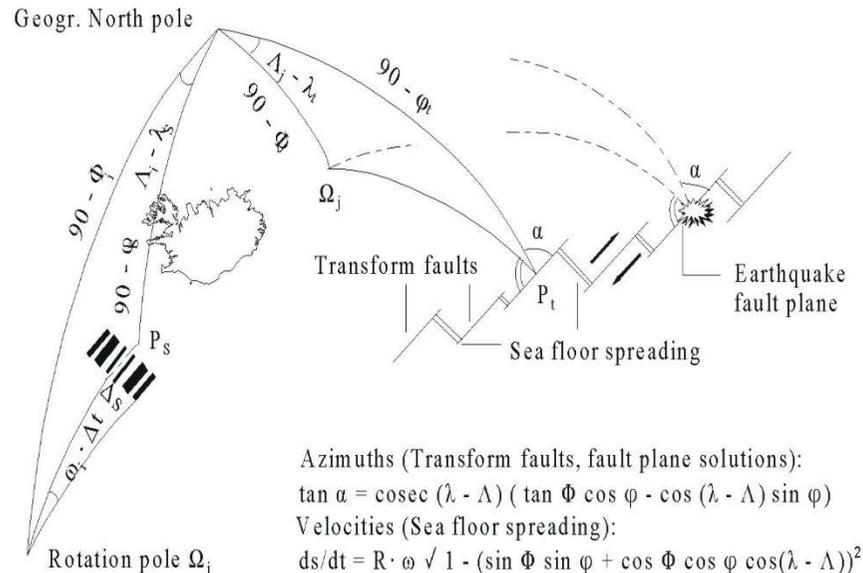
# Actual Continuous Kinematic Model (ACKIM) of the Earth's Crust based on ITRF2014

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# Introduction: Crustal kinematics based on plate tectonics

**Plate tectonic models** are based on geophysical data over geologic times.



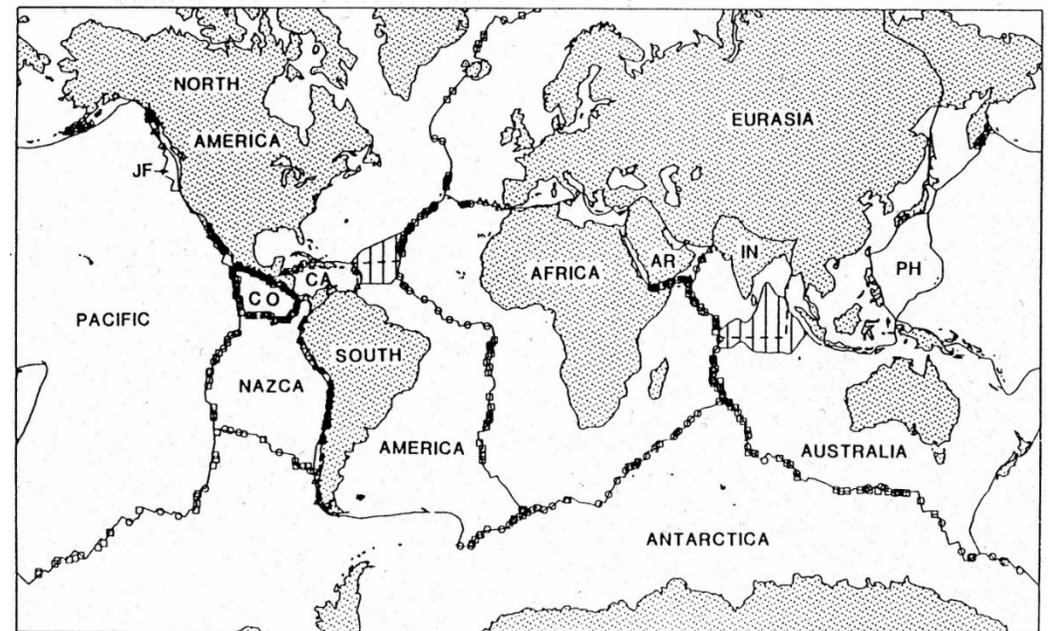
**Plate kinematics** is derived geophysically from three observation types:

- **Sea floor spreading** rates (velocities),
- **Transform faults** azimuths (directions),
- **Earthquake slip** vectors (directions).

The result are plate rotation vectors on a sphere (Theorem of Euler)

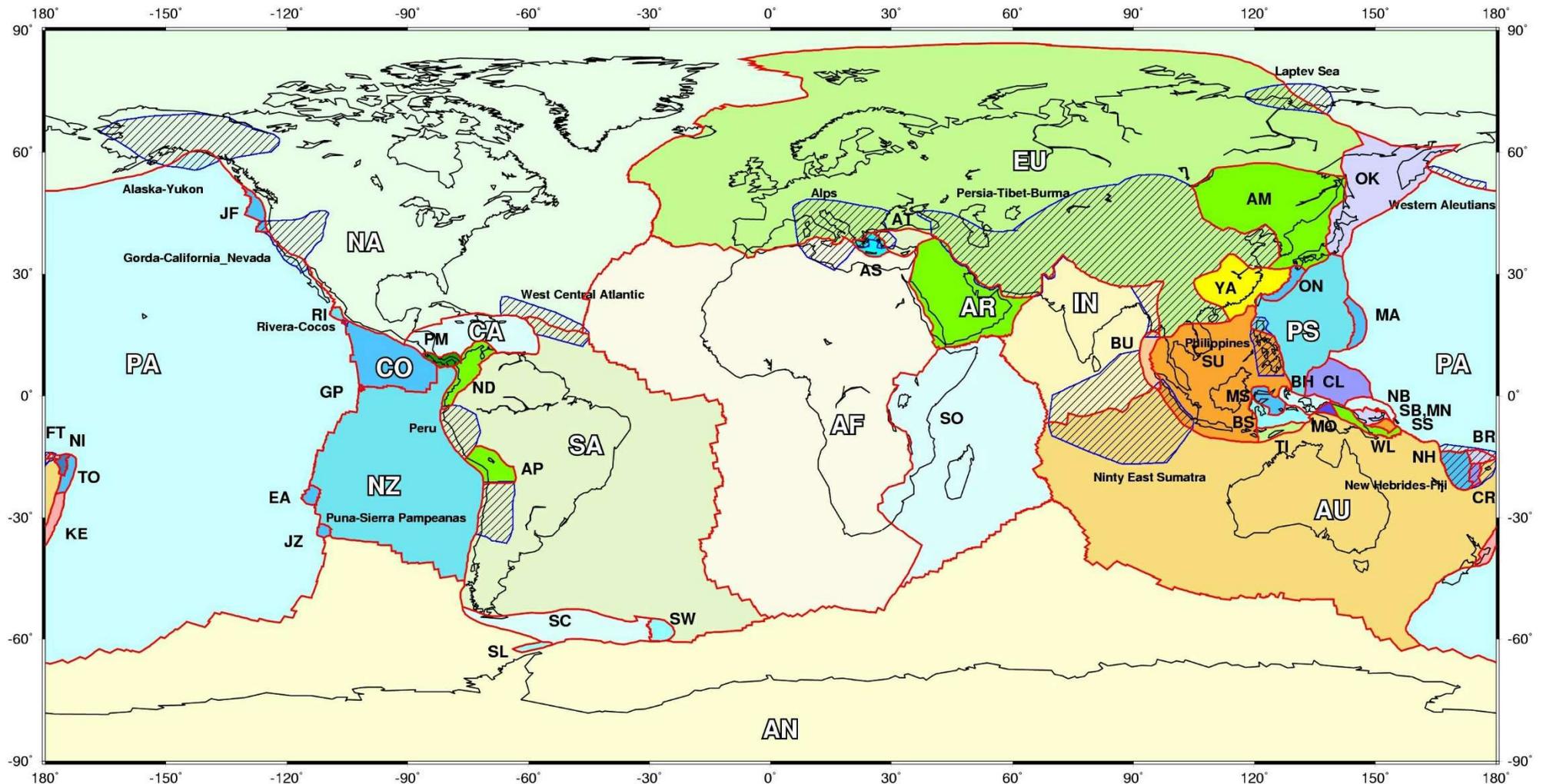
Such a model of rigid plates is NUVEL-1A (DeMets et al. 1994). Station velocities of the ITRF are based on this model defining the kinematic datum.

NUVEL-1A does not include any non-rigid crustal deformation.



# Considering more plates and crustal deformation zones

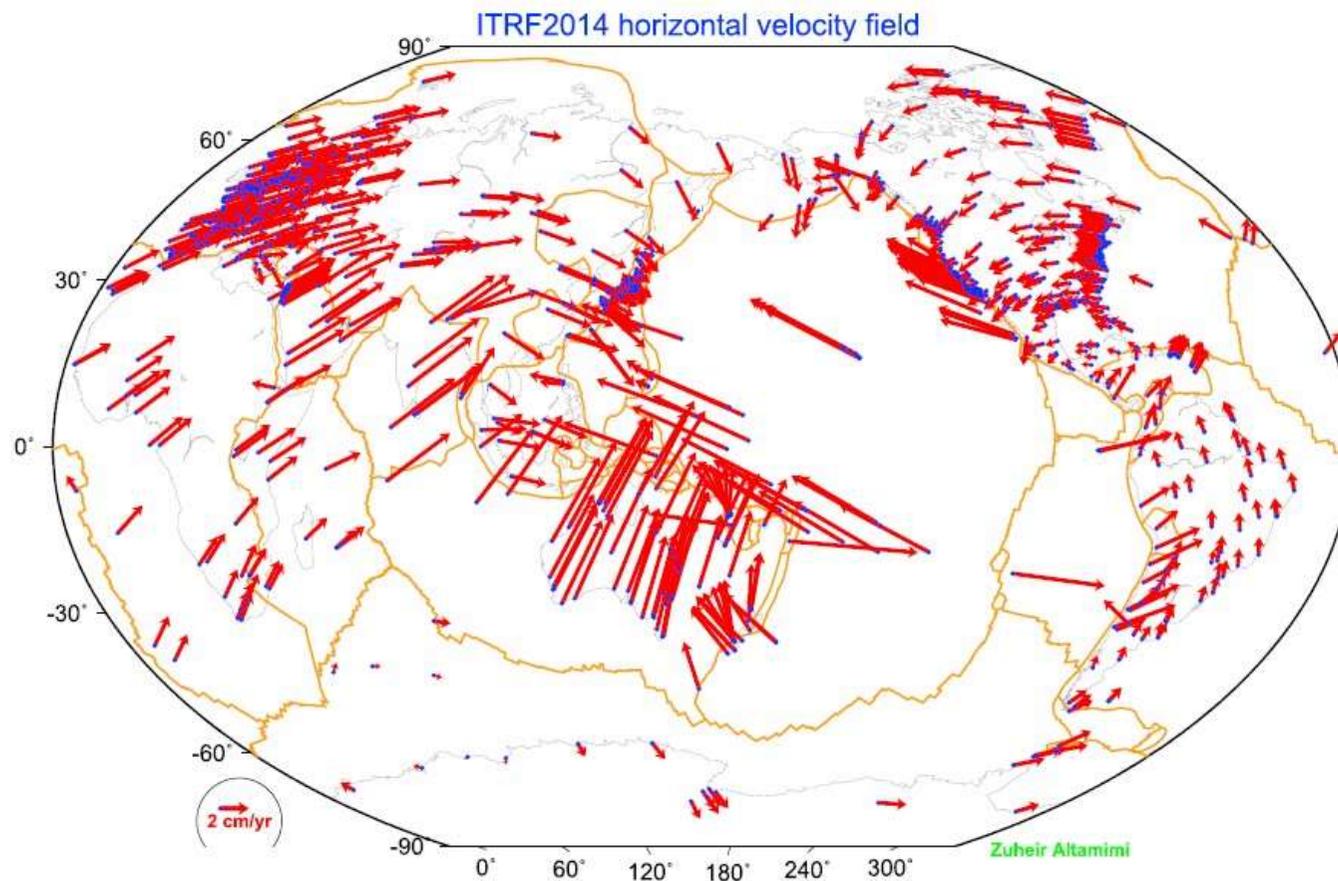
The plate model PB2002 (Bird 2003) includes 13 deformation zones within and between 52 plates, the 10 larger ones identical with NUVEL-1A.



The observation data cover a period of 3 Million years. Valid for today?

# Plate kinematics based on geodetic observations

- **Present day** plate kinematic models are only feasible since space geodetic observations allow **measuring global position changes** (velocities).
- Geodetic Actual Plate Kinematic Models (APKIM) are computed since 1988
- The latest APKIM2014 is based on the ITRF2014 (Altamimi et al. 2016)



## Velocities used

- only the latest periods in ITRF

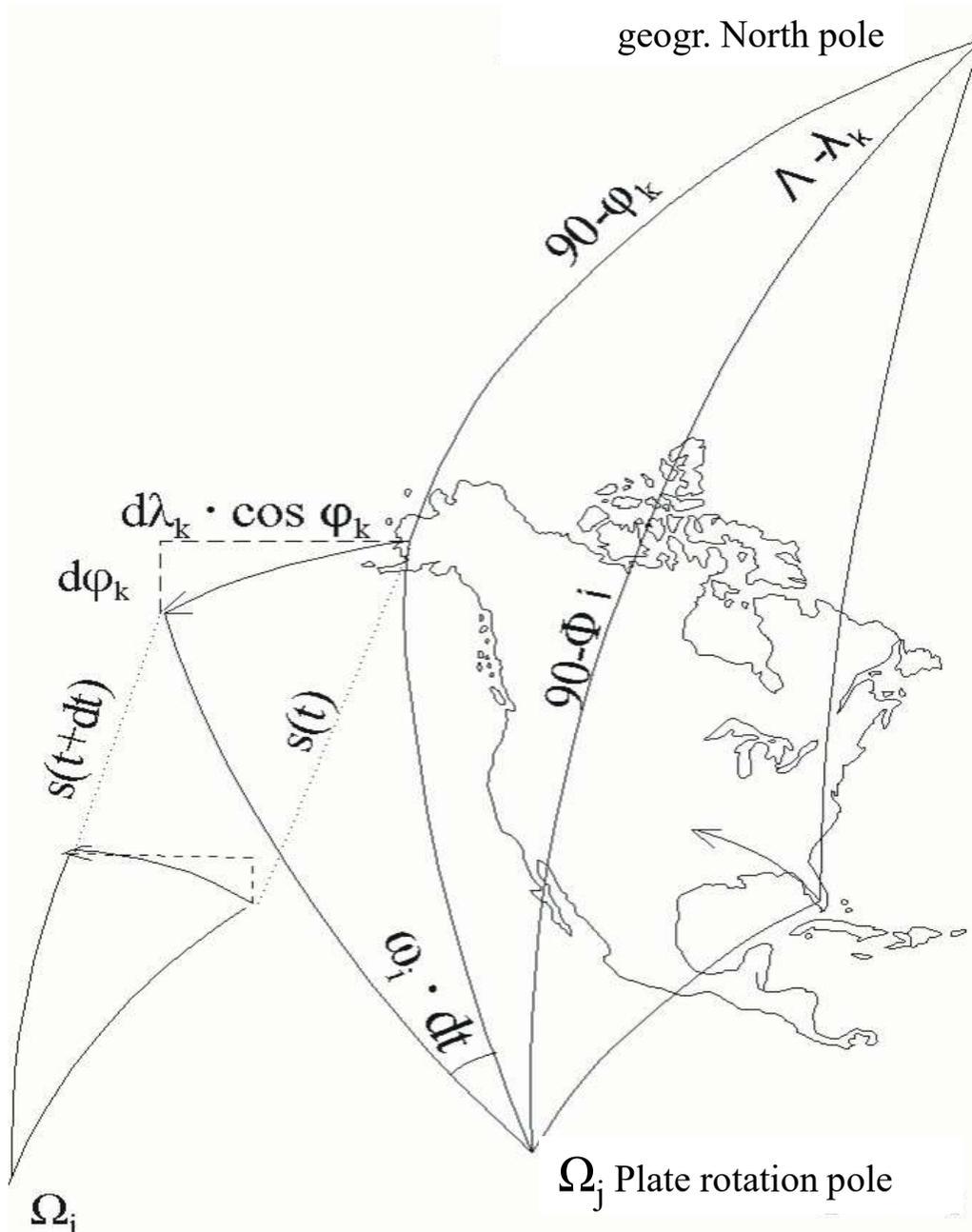
## Techniques

- GNSS 657
- Laser 46
- VLBI 54
- Doris 28

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Total 785

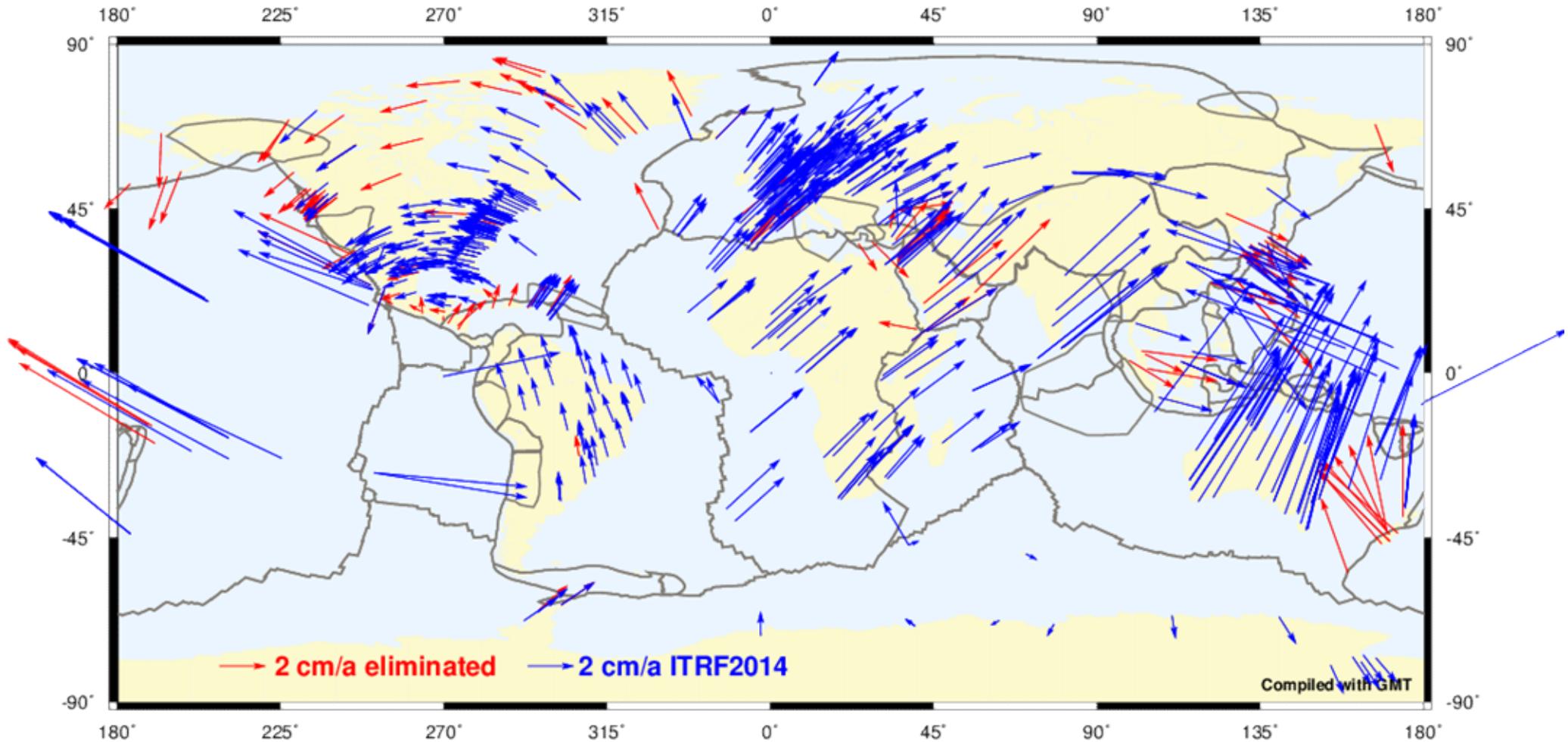
**Figure 11.** ITRF2014 horizontal site velocities with formal error less than 0.2 mm/yr. Major plate boundaries are shown according to *Bird* [2003].

# Summary of the processing procedure



- ITRF includes point coordinates and velocities of consecutive periods (solutions). A new period starts at any discontinuity.
- Only the latest periods are taken for the estimation of the plate rotation vectors ( $\Omega(\Phi, \Lambda, \omega)$ ).
- A two-dimensional adjustment is done (by spherical geometry) to avoid the effect of less precise vertical velocities.
- Iterative adjustments were done eliminating “non-fitting” ITRF velocities after the 3-sigma-criterion.

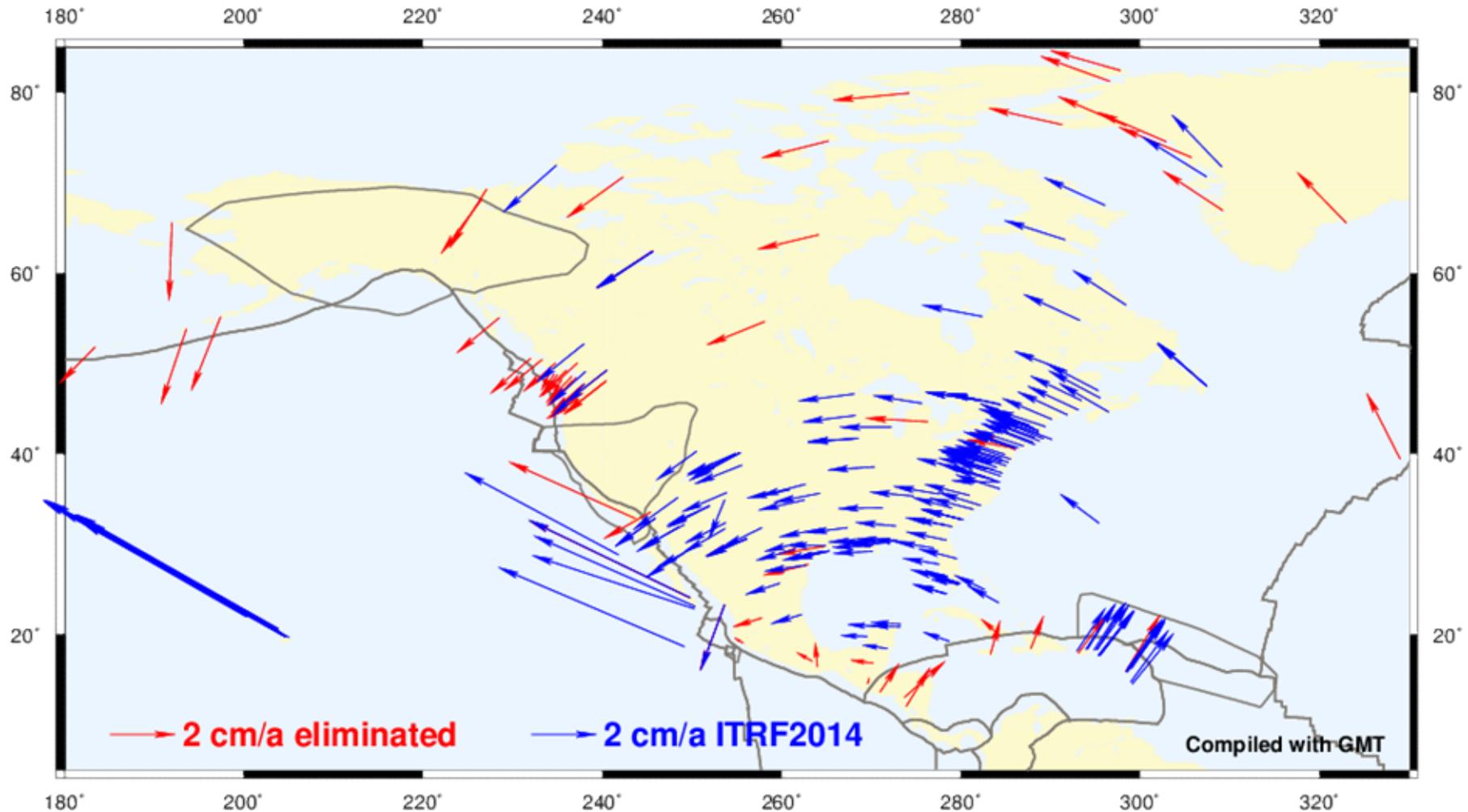
# Available & eliminated velocities (outliers / deformations)



ITRF2014 latest period: 785 velocity vectors; used: 636; eliminated: 149;  
Reasons: ITRF estimation uncertainties **or** intra-plate deformations

# Eliminated velocities (outliers / deformation-rigid plates)

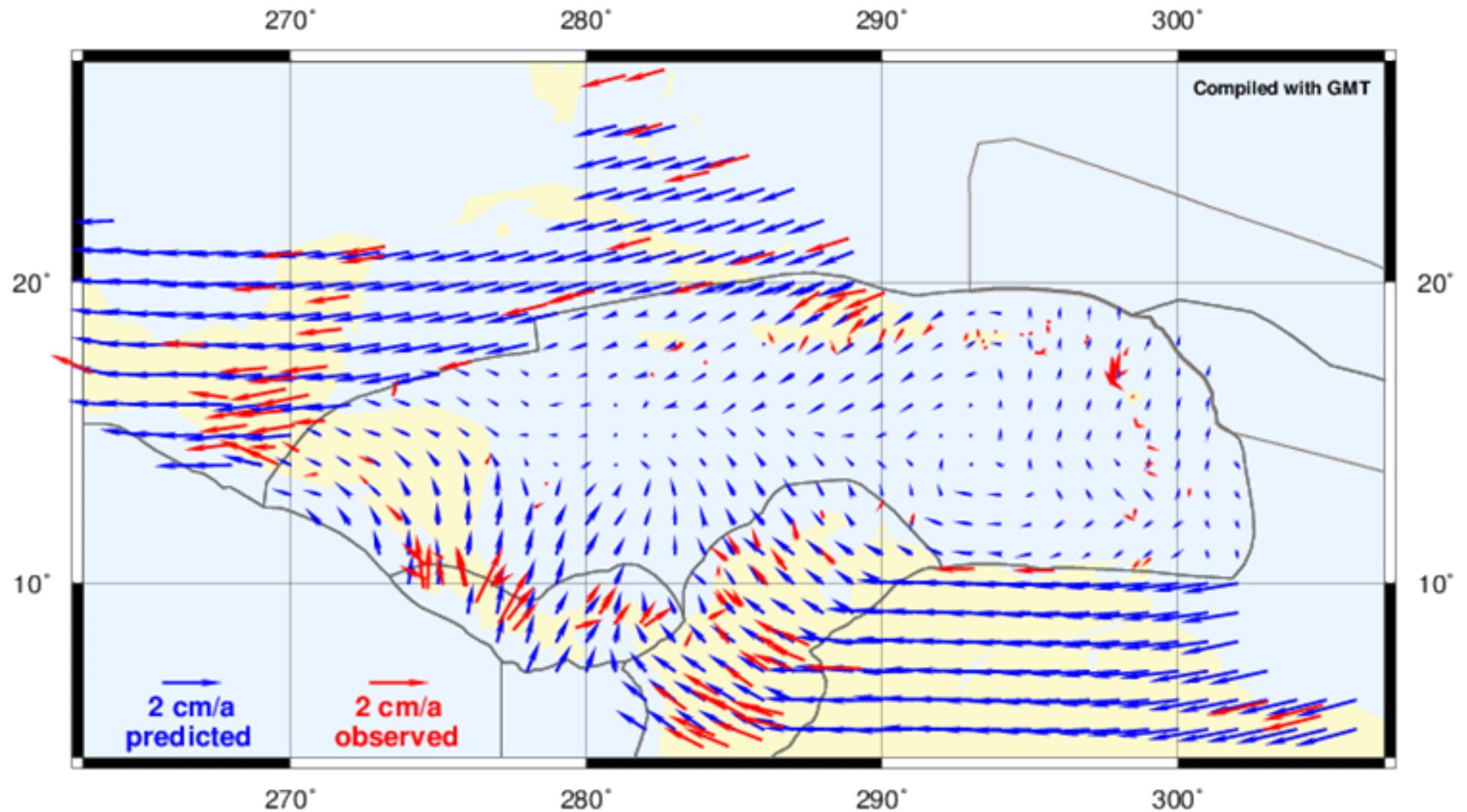
North American plate ITRF2014 available and eliminated velocity vectors



Most velocities are given in the southern part that dominates the estimation. Northern velocities do obviously not correspond to a rigid plate!

# Eliminated velocities (outliers / deformation-rigid plates)

Caribbean plate deformation (from the SIRGAS velocity model VEMOS2017)



The Caribbean plate is obviously not a rigid plate!

# Comparison of estimated plate rotation poles ( $\Phi$ , $\Lambda$ , $\omega$ )

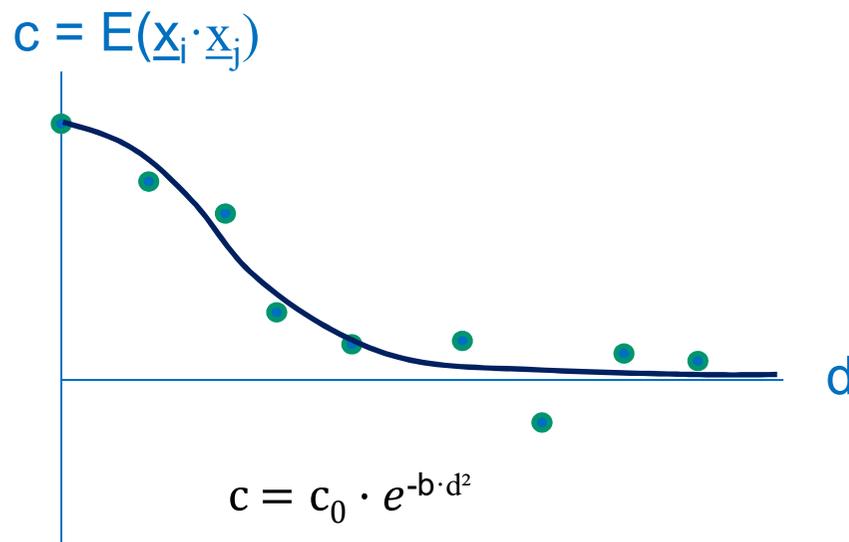
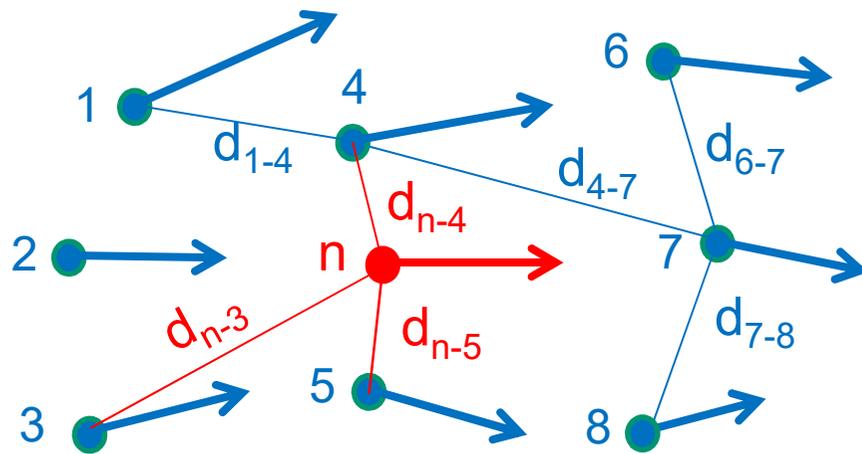


(red numbers are different to APKIM2014 after the 3-sigma-criterion)

Plate	APKIM2014			APKIM2008			NNR NUVEL-1A		
	$\Phi$ [°]	$\Lambda$ [°]	$\Omega$ [°/Ma]	$\Phi$ [°]	$\Lambda$ [°]	$\Omega$ [°/Ma]	$\Phi$ [°]	$\Lambda$ [°]	$\Omega$ [°/Ma]
Africa	49.57 ±0.19	278.71 ±0.54	0.267 ±0.001	49.80 ±0.26	278.54 ±0.70	0.268 ±0.001	50.57	286.04	0.291
Antarctica	59.32 ±0.39	234.04 ±0.56	0.216 ±0.003	58.83 ±0.33	231.91 ±0.59	0.214 ±0.003	62.99	244.24	0.238
Arabia	49.62 ±0.31	3.54 ±1.05	0.582 ±0.010	50.00 ±0.36	3.45 ±1.33	0.570 ±0.012	45.23	355.54	0.546
Australia	32.29 ±0.10	37.91 ±0.20	0.630 ±0.001	32.46 ±0.14	37.88 ±0.31	0.633 ±0.002	33.85	33.17	0.646
Caribbean	31.48 ±1.16	269.32 ±3.01	0.337 ±0.032	28.00 ±1.32	250.93 ±2.68	0.208 ±0.018	25.00	266.99	0.214
Eurasia	54.45 ±0.22	259.66 ±0.33	0.255 ±0.001	55.13 ±0.28	260.58 ±0.40	0.256 ±0.001	50.62	247.73	0.234
India	51.51 ±0.31	1.71 ±4.33	0.523 ±0.009	50.20 ±0.66	11.75 ±4.27	0.552 ±0.013	45.51	0.34	0.545
N. America	-4.82 ±0.30	272.10 ±0.13	0.193 ±0.001	-5.76 ±0.45	272.50 ±0.22	0.189 ±0.001	-2.43	274.10	0.207
Nazca	45.60 ±0.91	257.75 ±0.39	0.632 ±0.006	45.88 ±0.63	257.61 ±0.33	0.682 ±0.001	47.80	259.87	0.743
Pacific	-62.50 ±0.08	110.42 ±0.34	0.680 ±0.001	-62.57 ±0.08	110.93 ±0.36	0.634 ±0.005	-63.04	107.33	0.641
S. America	-18.68 ±0.51	231.31 ±1.30	0.122 ±0.001	-19.35 ±1.02	237.84 ±1.51	0.127 ±0.002	-25.35	235.58	0.116

# Alternative to plate models: continuous deformation model

An alternative to rigid plate kinematics for modeling global deformation is a continuous deformation model. We use a least squares collocation approach.



## 2D-vector prediction:

$$\underline{v}_{\text{pred}} = \underline{C}_{\text{new}}^T \underline{C}_{\text{obs}}^{-1} \underline{v}_{\text{obs}}$$

$\underline{v}_{\text{pred}}$  = predicted velocities ( $v_N, v_E$ )  
in a  $1^\circ \times 1^\circ$  grid

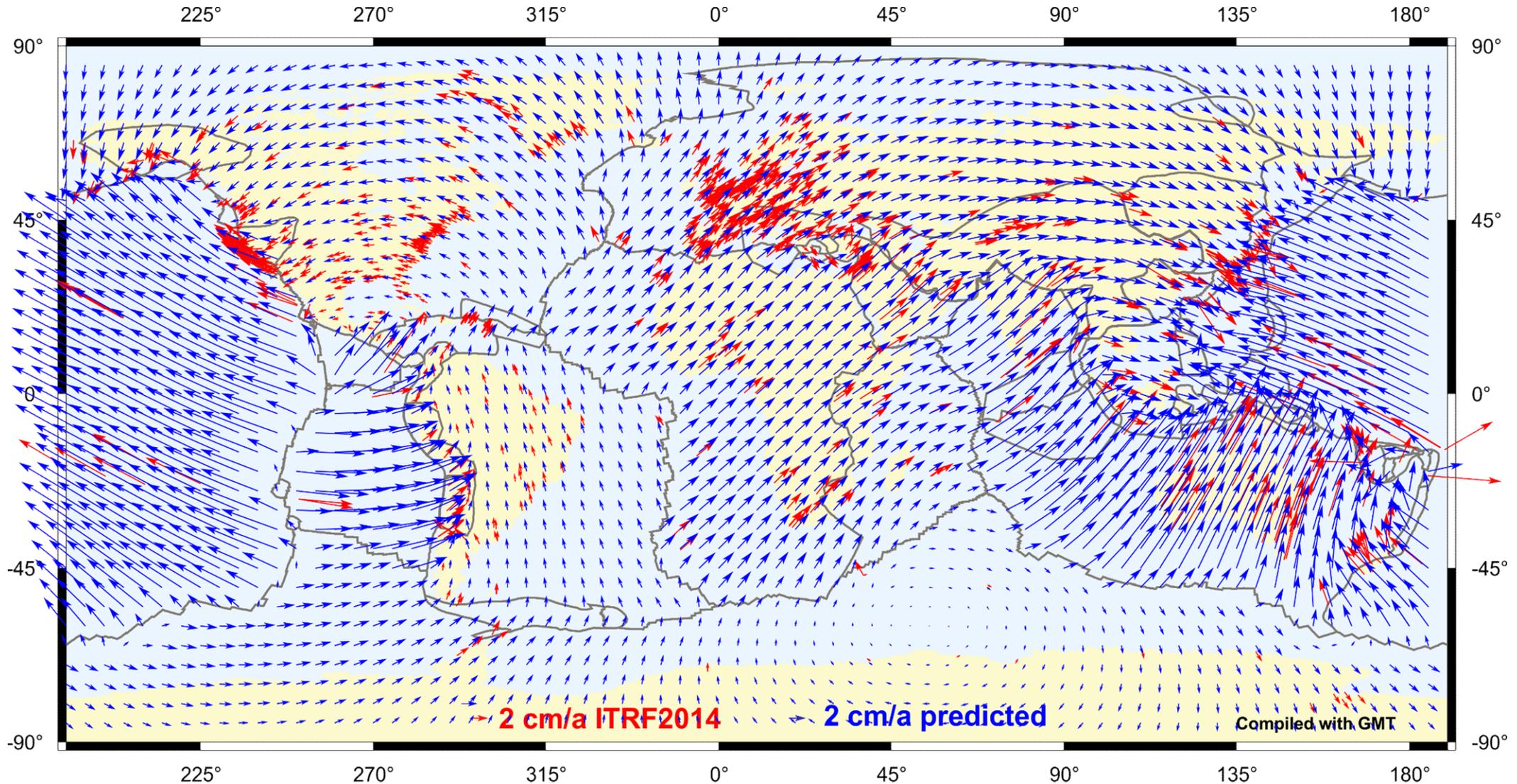
$\underline{v}_{\text{obs}}$  = observed velocities ( $v_N, v_E$ )  
in geodetic stations

$\underline{C}_{\text{new}}$  = correlation matrix between  
predicted & observed vectors

$\underline{C}_{\text{obs}}$  = correlation matrix between  
observed vectors ( $C_{NN}, C_{EE}, C_{NE}$ )

$\underline{C}$  matrices are built from empirical  
isotropic, stationary covariance  
functions  $c = E(\underline{x}_i \cdot \underline{x}_j)$ .

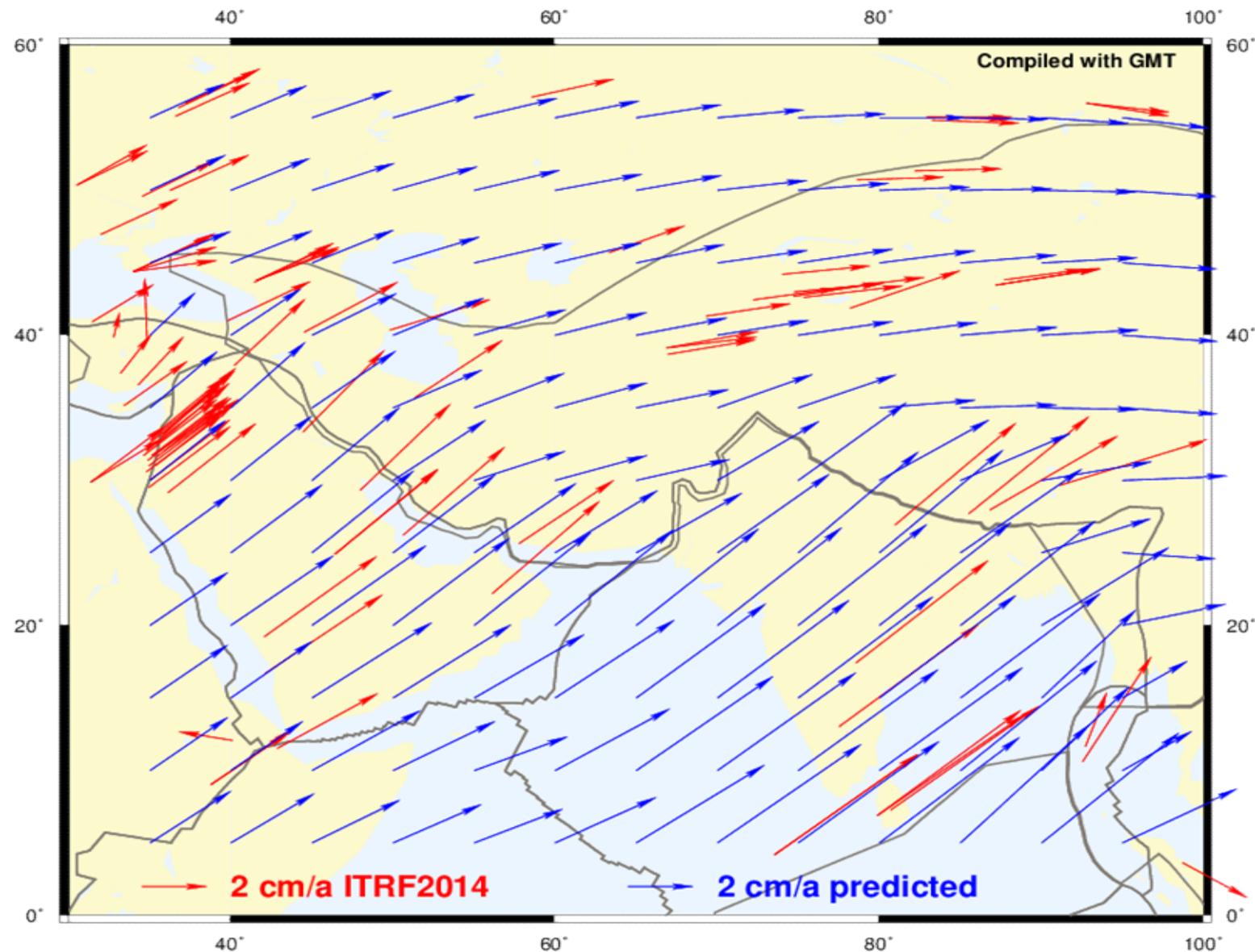
# Continuous crustal deformation model from ITRF2014



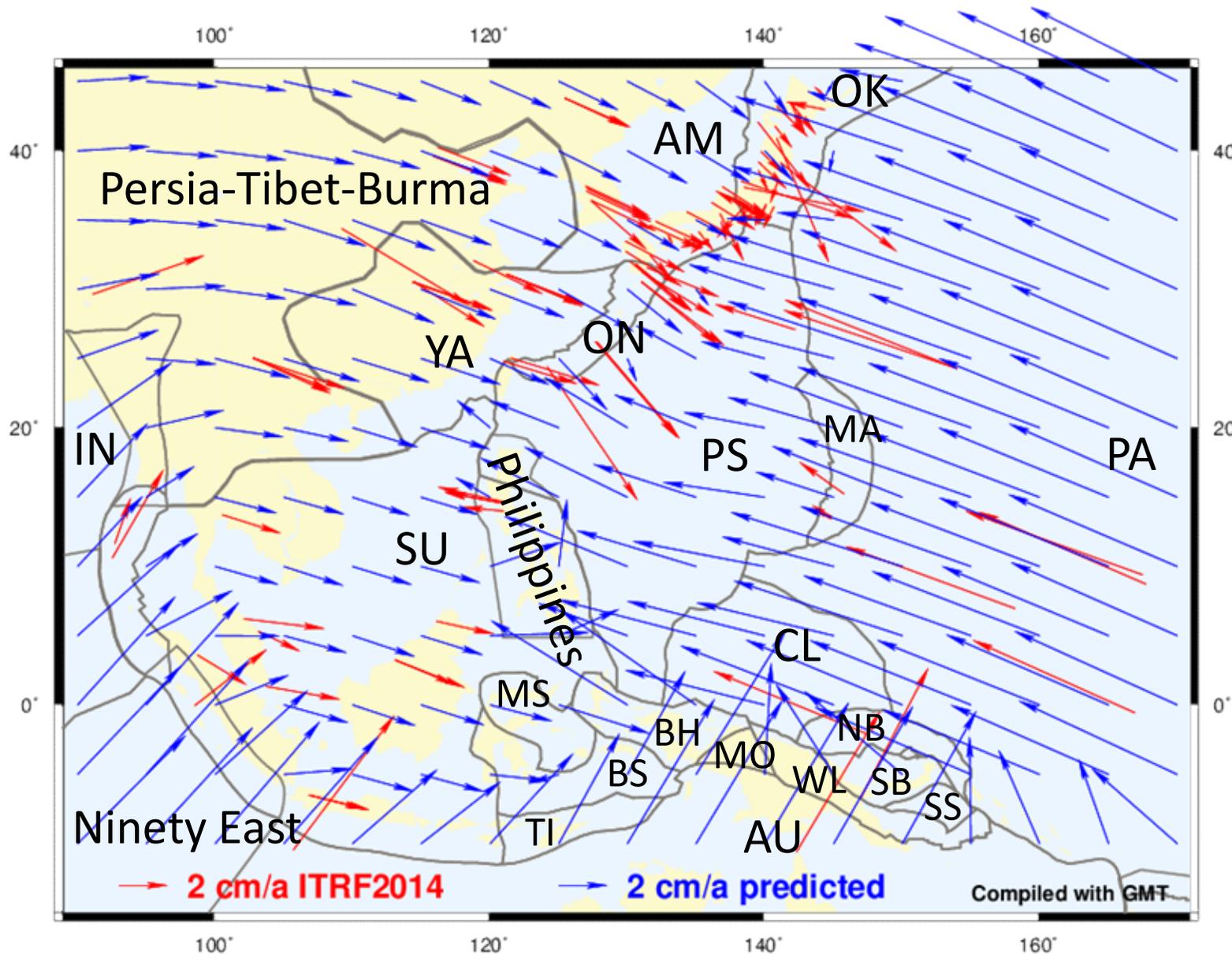
The collocation approach includes station velocities up to 500 km from the grid point to be predicted. If a sufficient number of station velocities ( $> 3$ ) is not available, the used plate model (APKIM2014) is considered.

# Example of continuous deformation zones

The Persia-Tibet-Burma orogen (Bird 2003) between Eurasia and India plates



# Example of extreme deformation zones



East Asia is a complicated tectonic area (extremely difficult for prediction)

Plates (e.g. PA) and orogens (deformation zone names) are according to Bird (2003)

- Plate tectonics is a good model for geology, in particular paleo geology.
- It can very well be used in geophysics for many modelling purposes.
- It may be used in geodesy as a basis for preliminary studies and models, but not for representing the global crustal deformation, which is needed in time-dependent global and continental reference frames (e.g. ITRF, AFREF, APREF, EUREF, NAREF, SIRGAS).

**Thank you very much for your attention!**