

Federal Agency for Cartography and Geodesy

SLR and the Gravity Field

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Gravity Field Determination

Dedicated gravity field satellite missions are available since several years



GRACE





GOCE

Why using SLR?

CHAMP

- Bad/unrealiable values of low-degree coefficients from GRACE: esp. C₂₀
- Gravity field variations priori to the launch of CHAMP, GRACE etc.
- Gravity field variations for the months with missing GRACE (K-Band) observations
- Filling the gaps if no dedicated gravity field mission is in operation, e.g., between GRACE and GRACE-FO

Gravity Field Determination and SLR

- SLR geodetic satellites (spherical satellites) orbit the Earth at higher altitudes than the satellite gravity missions (CHAMP, GRACE, GOCE etc.)
 - Lower sensitivity
- SLR observations are typically used for deriving
 - Low-degree gravity field coefficients (mainly degree 2)
 - Zonal harmonics
- Higher-degree monthly gravity field models can also be well derived from SLR observations using a combination of long and short arcs



GRACE vs. SLR



GRACE

Sensitive to ionosphere activity Active satellites, expensive maintenance Limited life-time

Normal points every 30 s (Starlette, Stella, AJISAI, LARES, Larets, BLITS) or every 120 s (LAGEOS) Quality of observations dependent on SLR stations (different frequencies and laser systems: 10Hz/kHz used) Strong correlations between some harmonics Directly connected to the terrestrial reference frame Different altitudes, typically above 800 km

SLR

No ionosphere delay of the signal Passive, low-cost satellites Unlimited life-time

SLR Satellites: Sensitivity to Gravity Field

| Perturbing accel. | Accel. on | Accel. on | Accel. on | Accel. on Stolla | |
|------------------------------------|---------------------|---------------------|---------------------|---------------------|--|
| Gravitational perturbations: | LAGEOD | AJISAI | LARES | Stella | |
| · Earth's monopole | 2.7 | 6.4 | 6.5 | 7.7 | |
| · Earth's oblateness C_{20} | $1.0\cdot 10^{-3}$ | $6.2\cdot 10^{-3}$ | $6.3\cdot 10^{-3}$ | $8.8\cdot 10^{-3}$ | |
| · Low-order grav. C_{22} | $6.0\cdot 10^{-6}$ | $3.6\cdot 10^{-5}$ | $3.7\cdot 10^{-5}$ | $5.1\cdot 10^{-5}$ | |
| \cdot Low-order grav. C_{66} | $8.6\cdot 10^{-8}$ | $3.1\cdot 10^{-6}$ | $3.2\cdot 10^{-6}$ | $6.3\cdot 10^{-6}$ | |
| \cdot Mid-order grav. C_{2020} | $8.1\cdot 10^{-13}$ | $1.5\cdot 10^{-8}$ | $1.6\cdot 10^{-8}$ | $1.1\cdot 10^{-7}$ | |
| \cdot Grav. attr. of Moon | $2.1 \cdot 10^{-6}$ | $1.4 \cdot 10^{-6}$ | $1.4 \cdot 10^{-6}$ | $1.3\cdot10^{-6}$ | |
| \cdot Grav. attr. of Sun | $9.6\cdot 10^{-7}$ | $6.4\cdot 10^{-7}$ | $6.5\cdot 10^{-7}$ | $5.7\cdot 10^{-7}$ | |
| \cdot Grav. attr. of Venus | $1.3\cdot 10^{-10}$ | $8.5\cdot 10^{-11}$ | $8.5\cdot 10^{-11}$ | $7.8\cdot 10^{-11}$ | |
| \cdot Solid Earth tides | $3.7\cdot 10^{-6}$ | $2.0\cdot 10^{-5}$ | $2.0\cdot 10^{-5}$ | $2.9\cdot 10^{-5}$ | |
| \cdot Ocean tides | $3.7\cdot 10^{-7}$ | $1.9\cdot 10^{-6}$ | $2.0\cdot 10^{-6}$ | $3.0\cdot 10^{-6}$ | |
| General relativity: | | | | | |
| \cdot Schwarzschild effect | $2.8\cdot 10^{-9}$ | $1.1\cdot 10^{-8}$ | $1.1\cdot 10^{-8}$ | $1.4\cdot 10^{-8}$ | |
| \cdot Lense-Thirring effect | $2.7\cdot 10^{-11}$ | $1.3\cdot 10^{-10}$ | $1.4\cdot 10^{-10}$ | $1.8\cdot 10^{-10}$ | |
| \cdot Geodetic precession | $3.4\cdot 10^{-11}$ | $4.2\cdot 10^{-11}$ | $4.2\cdot 10^{-11}$ | $4.3\cdot 10^{-11}$ | |
| Non-gravitational perturbations: | | | | | |
| \cdot Solar radiation pressure | $3.5\cdot 10^{-9}$ | $2.5\cdot 10^{-8}$ | $1.1\cdot 10^{-9}$ | $4.4\cdot 10^{-9}$ | |
| \cdot Earth radiation pressure | $4.4\cdot 10^{-10}$ | $8.6\cdot 10^{-9}$ | $3.9\cdot 10^{-10}$ | $1.8\cdot 10^{-9}$ | |
| \cdot Thermal re-radiation | $5.0\cdot 10^{-11}$ | $4.1\cdot 10^{-10}$ | $1.9\cdot 10^{-11}$ | $6.9\cdot 10^{-11}$ | |
| \cdot Light aberration | $1.1\cdot 10^{-13}$ | $1.1\cdot 10^{-12}$ | $5.1\cdot 10^{-14}$ | $2.0\cdot 10^{-13}$ | |
| \cdot Atmospheric drag (~ min) | $0.8\cdot 10^{-14}$ | $3.0\cdot 10^{-11}$ | $2.6\cdot 10^{-12}$ | $5.0\cdot 10^{-11}$ | |
| \cdot Atmospheric drag (~ max) | $2.0\cdot 10^{-13}$ | $5.9\cdot10^{-10}$ | $4.8\cdot 10^{-11}$ | $5.0\cdot 10^{-8}$ | |

Sensitivity shows clear dependency on orbital height

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SLR Satellites: Sensitivity to Gravity Field



Current ILRS standard setup

BUT: Satellites with lower altitudes are needed for gravity field determination

SLR Satellites: Sensitivity to Gravity Field

With an increased number of used satellites, the solution gains sensitivity w.r.t. the Earth's gravity field:

- 4 satellites: up to d/o 3
- 5 satellites: up to d/o 6 (higher d/o for tesseral coefficients)
- Max. constellation: up to d/o 12

But:

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Not all coefficients can be determined reliably due to remaining correlations!



SLR Satellites: Correlations between Gravity Field Coefficients



- The correlation between C_{2,0} and C_{3,0} can be reduced significantly by using additional satellites.
- In an 11-satellite constellation, both parameters are decorrelated.
- The correlation between $C_{2,0}$ and $C_{4,0}$ cannot be eliminated.
- Reason: there is no satellite orbit sensitive to only one of these coefficients (geometrical correlation of both coefficients).

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SLR gravity solutions: Parameter set-up

| | | SLR solutions | | |
|------------------------------|-----------------------------|---|-------------------|--|
| Esti | imated parameters | LAGEOS-1/2, Starlette, Stella, AJISAI, LARES, Blits, Larets, Beacon-C | H | |
| | Osculating elements | a, e, i, Ω, ω, u ₀ (LAGEOS: 1 set per 10 days, LEO: 1 set per 1 day) | So So Al | |
| Orbits | Dynamical parameters | $\begin{array}{c} LAGEOS-1/2:S_0,S_S,S_C\\ (1 \text{ set per 10 days})\\ Sta/Ste/AJI:C_D,S_C,S_S,W_C,W_S\\ (1 \text{ set per day}) \end{array}$ | Bo | |
| | Pseudo-stochastic pulses | LAGEOS-1/2 : no pulses Sta/Ste/AJI : once-per-revolution in along-track only | av D | |
| Earth rotation parameters | | X _P , Y _P , UT1-UTC (Piecewise linear, 1 set per day) | ur us | |
| Geocenter coordinates | | 1 set per 30 days | | |
| Earth gravity field | | Estimated up to d/o 10/10 (1 set per 30 days) | | |
| Station coordinates | | 1 set per 30 days | | |
| 0 | ther parameters | Range biases for all stations (LEO) and for selected stations (LAGEOS) | ty Field I SIRGAS | |

Here: Set-up of the solutions from AIUB using the Bernese GNSS Software

Similar solutions available from DGFI-TUM with up to 11 satellites using DOGS

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SLR gravity solutions (Bernese GNSS Software)



Selection of satellites: Correlations between parameters

Correlation matrix of combined solution comprising orbit parameters (LA-1 only), GFC and EOP



- orbital elements and p_{albe} (a)
- Stokes coefficients (b)
- (c) EOP and SRP scaling factor (p_{rad})
- (d), (e), (f) correlations between parameter groups





Selection of satellites: Correlations between parameters

Correlation matrix of combined solution comprising orbit parameters (LA-1 only), GFC and EOP



- (a) orbital elements and $p_{\rm albe}$
- (b) Stokes coefficients
- (c) EOP and SRP scaling factor (p_{rad})
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Selection of satellites: Correlations between parameters

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SLR gravity solutions: Comparison with GRACE Annual signal



Associated cumulative distribution function showing the significance of the recovered annual signal for SLR solutions (left) and GRACE solutions (right). => SLR mostly sensitive to d/o 6/6

SLR gravity solutions: Comparison with GRACE Secular changes

- Secular changes of geoid deformations derived from SLR show a very high level of consistency with the GRACEbased results, however, with a lower spatial resolution.
- The ice mass loss in Greenland, West Antarctica and Patagonia is well captured in the SLR solutions.

Reference: Sośnica, K., Jäggi, A., Meyer, U., Thaller, D., Beutler G., Arnold, D., Dach, R. (2015). *Time variable Earth's gravity field from SLR satellites*. Journal of Geodesy, on-line (http://link.springer.com/article/10.1007/s00190 -015-0825-1)



SLR gravity solutions: Comparison with GRACE Mass change in individual regions





The SLR solutions can recover the largest seasonal and secular variations of the gravity field, which correspond to the largescale mass transport in the system Earth, e.g., the accelerating ice mass depletion in Greenland.

The amplitudes in the SLR solutions up to d/o 10/10 are typically underestimated due to the limited sensitivity of SLR solutions to coefficients of degree 7-10.

SLR + GRACE (GPS) + GRACE (K-Band)

Can SLR contribute anything at all?

SLR contributes most to the zonal gravity field coefficients and the coefficients of degree-2.

 C_{20} is degraded in the GRACE solutions due to long-period signals, because some signals in C_{20} have the same period as the S₂ and S₁ tidal aliases with GRACE orbits. This degradation is reflected in a large peak of 160 days in the C₂₀ spectra.

Combination of GRACE with SLR solutions remarkably reduces the spurious 160-day peak in C₂₀ series.





Spectra of monthly C20 -values from 2003-2013. GRACE solutions are affected by spurious signal at 160d-period that is cured by the combination with SLR.

Gravity Field and Orbit Determination

 C_{20} causes an acceleration:

$$\begin{pmatrix} R \\ S \\ W \end{pmatrix} = \frac{3}{2} \frac{GMa_e^2 \Delta C_{20}}{r^4} \begin{cases} 1 - \frac{3}{2} \sin^2 i + \frac{3}{2} \sin^2 i \sin^2 u \\ \sin^2 i \sin 2u \\ \sin 2i \sin u \end{cases}$$

(in Radial *R*, along-track *S*, out-of-plane *W*)

Correlation between a change in C₂₀ and a once-per-Rev acceleration in the outof-plane direction *W*

Impact on orbit parameterization !



Gravity Field and Orbit Determination

 C_{20} causes an acceleration:

$$C_{20} \text{ causes an acceleration:} \quad \begin{cases} R \\ S \\ W \end{cases} = \frac{3}{2} \frac{GMa_e^2 \Delta C_{20}}{r^4} \begin{cases} 1 - \frac{3}{2} \sin^2 i + \frac{3}{2} \sin^2 i \sin^2 u \\ \sin^2 i \sin 2u \\ \sin 2i \sin u \end{cases}$$
(in Radial *R*, along-track *S*, out-of-plane *W*)



Figure 3.2: Empirical out-of-plane W_S and W_C parameters for LAGEOS-1 (left) and LAGEOS-2 (right). Note the different scales.

Gravity Field and Earth Rotation: LOD





Federal Agency for Cartography and Geodesy Simultaneous estimation of LOD and Gravity Field is needed to:

- Reduce the offset of LOD estimates
- Reduce the a posteriori error of LOD estimates
- Reduce artificial signals with periods corresponding to orbit modelling issues (e.g., draconitic year)

Gravity Field and Earth Rotation: LOD



Systematic error in LOD deflects the dUT polygon

A combination with different SLR satellites reduces the systematics.

- Decrease of the systematic LOD error:
- Factor 4 due to inclusion of LARES (in addition to LAG-1/-2)
- Factor 10 due to the 10-satellite combination

Thank you for your kind attention!

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Consistent dynamic satellite reference frames and terrestrial geodetic datum parameters

• Due to the high sensitivity of SLR observations to the fundamental geodetic parameters, correlations might falsify reliable estimates



Correlations related to Stokes coefficients

b) correlations of orbit parameters and Stokes coefficients

c) correlations of LOD and C_{l0} ; x_p/y_p with C_{21}/S_{21}

e) correlation of TRF scale with C_{00} ; origin with $C_{10}/C_{11}/S_{11}$; orientation Federal Agency for Cartography and Geodes with $C_{21}/S_{21}/C_{22}/S_{22}$

+ Improvement- Degradation

Summary

滃

| | | 4-sat. | 11-s | at. | 4-sat | + AJI | 4-sat. + | STA | 4-sat. + | STE | 4-sat. + | LTS | 4-sat. + | LRS |
|---------------|--|------------------|--------|--------------|------------|---------------|----------|------|----------|---------------|----------|---------------|------------|--------------|
| | | (ref.) | value | [%] | value | [%] | value | [%] | value | [%] | value | [%] | value | [%] |
| nm] | WRMS (Tx) | 6.2 | 5.0 | 11.3 | 6.0 | 3.2 | 5.4 | 12.9 | 5.3 | 14.5 | 5.5 | 11.3 | 5.3 | 14.5 |
| | wmean (Tx) | 3.3 | 2.5 | 24.2 | 1.6 | 51.5 | 2.8 | 15.2 | 3.5 | 6.1 | 3.0 | 9.1 | 3.3 | 0.0 |
| | | | | | | | | | | | | | | |
| | WRMS (Ty) | 6.3 | 6.0 | 4.8 | 6.3 2.9 | 20.4 | 5.7 | 9.5 | 5.6 | 11.1 | 5.7 | 9.5 | 5.5 2.7 | 12.7 |
| | winean (1y) | -4.0 | -2.9 | 57.0 | -3.2 | 30.4 | -3.3 | 20.3 | -3.7 | 19.0 | -4.0 | 0.0 | -3.1 | 19.0 |
| | WRMS (Tz) | 9.8 | 6.4 | 34.7 | 7.5 | 23.5 | 9.5 | 3.1 | 8.1 | 17.3 | 9.5 | 3.1 | 9.0 | 8.2 |
| E E | wmean (Tz) | 0.5 | -2.6 | 420.0 | -3.0 | 500.0 | -1.2 | 140 | -0.2 | 60.0 | -0.3 | 40.0 | 0.0 | 100.0 |
| \mathbf{TR} | | | | | | | | | | | | | | |
| | WRMS (Sc) | 6.7 | 5.4 | 19.4 | 7.6 | 13.4 | 5.7 | 14.9 | 6.7 | 0.0 | 6.6 | 1.5 | 5.9 | 11.9 |
| | wmean (Sc) | 4.7 | 9.3 | 97.9 | 16.0 | 240.4 | 5.0 | 6.4 | 4.5 | 4.3 | 5.2 | 10.6 | 4.3 | 8.5 |
| | WRMS (north) | 11.9 | 9.3 | 21.8 | 11.3 | 5.0 | 10.5 | 11.8 | 10.8 | 9.2 | 11.1 | 6.7 | 11.1 | 6.7 |
| | WRMS (east) | 10.8 | 8.7 | 19.4 | 10.0 | 7.4 | 10.1 | 6.5 | 9.9 | 8.3 | 10.6 | 1.9 | 10.6 | 1.9 |
| | WRMS (height) | 12.7 | 11.8 | 7.1 | 12.1 | 4.7 | 12.5 | 1.6 | 12.6 | 0.8 | 12.3 | 3.1 | 12.6 | 0.8 |
| | WRMS (x-pole) | 0.385 | 0.320 | 16.9 | 0.360 | 6.5 | 0.356 | 7.5 | 0.366 | 4.9 | 0.361 | 6.2 | 0.414 | 7.5 |
| [p] | wmean $(x-pole)$ | -0.107 | 0.004 | 96.3 | -0.053 | 50.5 | -0.052 | 51.4 | -0.079 | 26.2 | -0.069 | 35.5 | -0.102 | 4.7 |
| ms | | | | | | | | | | | | | | |
| as, | WRMS $(y-\text{pole})$ | 0.387 | 0.305 | 21.2 | 0.355 | 8.3 | 0.352 | 9.0 | 0.353 | 8.8 | 0.355 | 8.3 | 0.384 | 0.8 |
| <u></u> | wmean $(y-pole)$ | 0.165 | 0.027 | 83.6 | 0.035 | 78.8 | 0.099 | 40.0 | 0.118 | 28.5 | 0.134 | 18.8 | 0.193 | 17.0 |
| OP | WDMS (ALOD) | 0.047 | 0.042 | 0 5 | 0.040 | 14.0 | 0.020 | 17.0 | 0.026 | 02.4 | 0.025 | 05.5 | 0.029 | 10.1 |
| E | wran (ΔLOD) | -0.047 | 0.043 | 8.8 200.0 | -0.006 | 14.9 500.0 | 0.039 | 800 | 0.036 | 23.4 100.0 | 0.035 | 20.0 300.0 | 0.038 | 19.1 |
| _ | | -0.001 | 0.005 | 200.0 | -0.000 | 000.0 | 0.005 | 000 | 0.002 | 100.0 | 0.004 | 300.0 | 0.001 | 0.0 |
| - | WRMS $(C_{(2,0)})$ | 2.0172 | 1.0554 | 47.7 | 1.1151 | 44.7 | 1.2835 | 36.4 | 1.1945 | 40.8 | 1.2655 | 37.3 | 1.1813 | 41.4 |
| 1-1 | WRMS $(C_{(3,0)})$ WPMS $(C_{(3,0)})$ | 17.224 6.0205 | 1.2346 | 92.8 87.9 | 0.9831 | 94.3 42.7 | 14.726 | 14.5 | 8.4124 | 51.2 84.5 | 10.414 | 39.5 | 7.7003 | 55.3 81.0 |
| ö | $(U_{(4,0)})$ | 0.9300 | 0.0001 | 01.2 | 0.0990 | 40.1 | 4.0411 | 41.1 | 1.0100 | 04.0 | 1.1040 | 14.0 | 1.0144 | 51.0 |
| GF | WRMS* $(C_{(n,m)})$ | | | 79.3 | | 64.4 | | 66.1 | | 52.3 | | 41.9 | | 61.7 |

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Estimation of reliable low degree spherical harmonics



Solution setup as used by CSR/GSFC: Estimation of a reliable subset of coefficients



Scatter of the centered degree-1 Stokes coefficient solutions





- CSR SLR RL05: 5-satellite solution (Cheng et al., 2013)
- AIUB SLR: 8-satellite solution (Sosnica et al., 2015)
- DGFI-TUM SLR: 11-satellite solution, TRF and EOP fixed (Bloßfeld et al., submitted)
- Swenson (2008): GRACE, geophysical model for ocean bottom pressure (OBP)
- Rietbroek (2016): GRACE, OBP, GPS

Geocenter: Single techniques

Draconitic year is visible: GNSS = 352d, LAG-2 = 222d



Validation of low degree spherical harmonics (M. Talpe)

How can we validate the SLR-only low degree spherical harmonics?

- Compare seasonal signals of individual coefficients
- Compare trend signals of individual coefficients
- Map-based comparison with GRACE since individual coefficients are correlated
- Investigation of mass change in Greenland and Antarctica by comparison with GRACE
- Compare annual/semi-annual signals of maps (not shown here)
- Compare degree variances of solutions (Bloßfeld et al., 2015)
- Compare EOF modes to capture the main signal content (not shown here)
- Investigate spatial coherence (not shown here)



- How can we validate the SLR-only low degree spherical harmonics?
 - Compare seasonal signals of individual coefficients (M. Talpe)



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- How can we validate the SLR-only low degree spherical harmonics?
 - Compare trend signals of individual coefficients (M. Talpe)



- How can we validate the SLR-only low degree spherical harmonics?
 - Map-based comparison with GRACE since individual coefficients are correlated (M. Talpe)
 - All solutions are expanded in Equivalent Water Thickness maps and compared to their GRACE equivalent. The RMS of the difference of the two



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• How can we validate the SLR-only low degree spherical harmonics?



- How can we validate the SLR-only low degree spherical harmonics?
 - Investigation of mass change in Greenland and Antarctica by comparison with GRACE (M Talpe)

